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| Title:        | Seismic Analysis of the 12 August 2000 Kursk Submarine Disaster in the Barents Sea |
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## Seismic Analysis of the 12 August 2000 Kursk Submarine Disaster in the Barents Sea

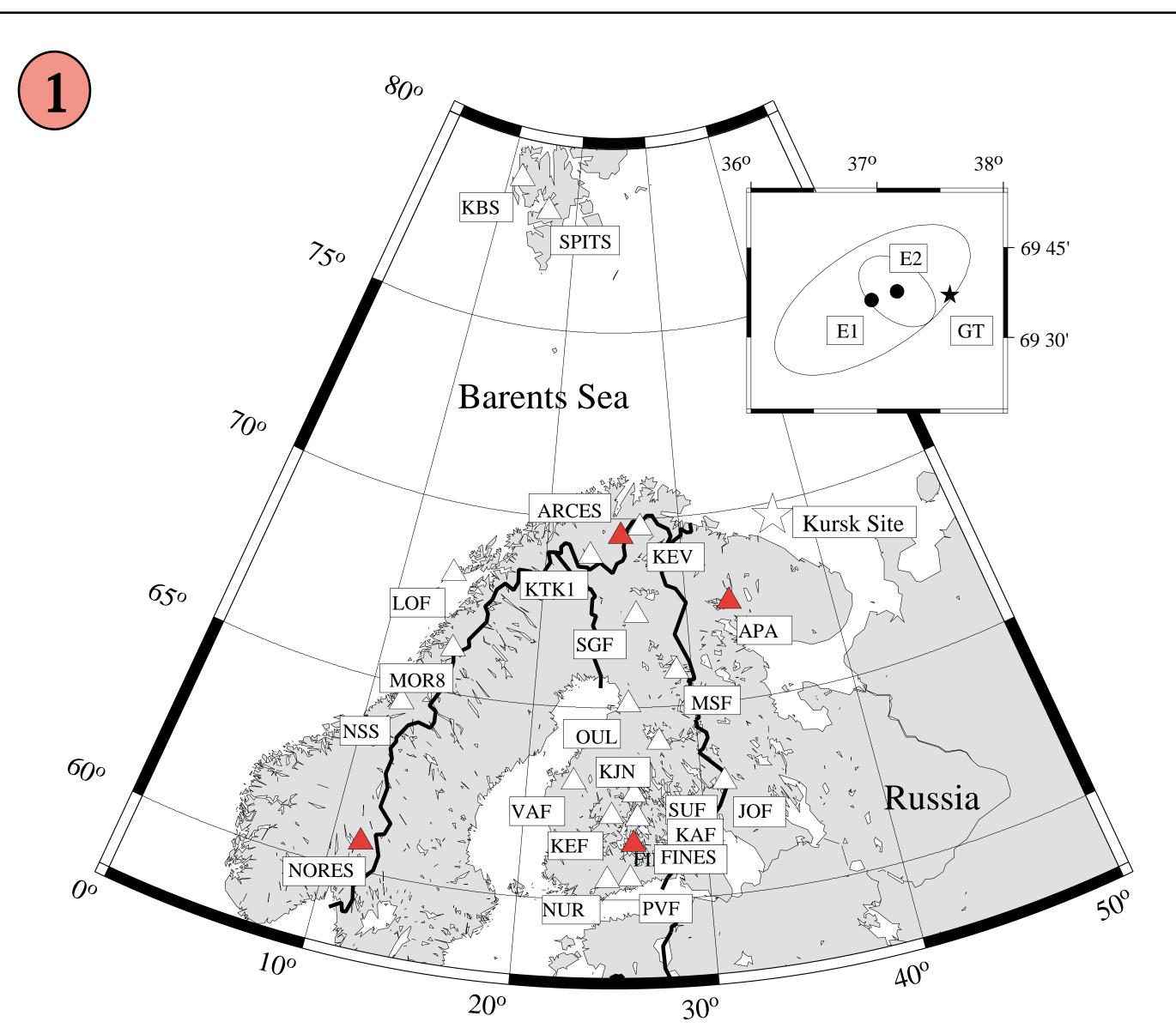


Figure 1. Locations of regional distance seismic stations which recorded the Kursk event. Three component stations are shown as white triangles and short period array stations are shown as red triangles. The inset shows our locations and error ellipses for the precursory event (E1) and the main event (E2). The actual position (GT) of the downed sub is indicated by the star.

Second



Introduction

On 12 August 2000 the most advanced attack submarine in the Russian fleet sank in the Barents Sea about 70 km off the coast of the Kola Peninsula. Details of the accident that caused the Kursk to sink were shrouded in secrecy and propaganda. It is known that the accident was related to a series of explosions, and owing to the high efficiency of coupling in underwater explosions these explosions generated a seismic signal of sufficient magnitude (4.2 M<sub>1</sub>, 3.4 m<sub>b</sub>) to be recorded by seismometers as far as 5000km away. The main Kursk event on August 12 was detected and automatically located by at least four independent monitoring groups. The preliminary locations are:

> Univ. Helsinki: 69.67°N 37.53°E 07:30:41.9 69.58°N 38.03°E 07:30:42.6 69.67°N 37.25°E 07:30:42.0 NORSAR: PIDC: 69.58°N 37.92°E 07:30:42.2

The differences in location, which are negligible, are due to the different techniques, models and data sets. The seismic signal can be positively associated with the Kursk incident due to origin time and lack of previous seismicity in the region.



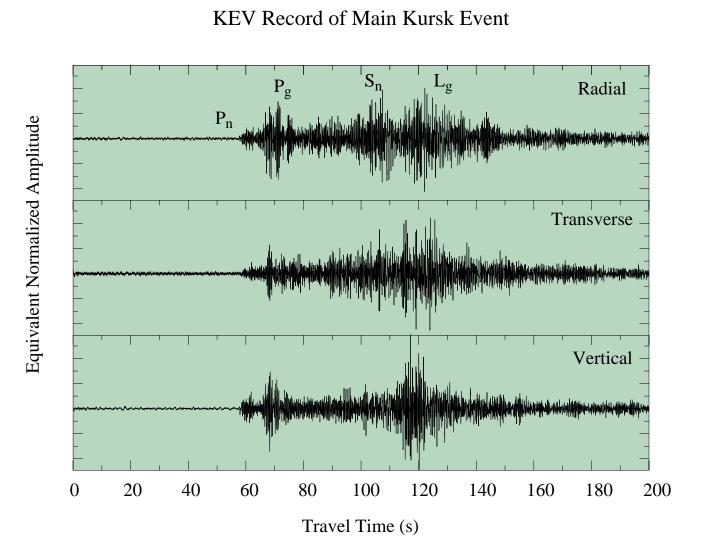
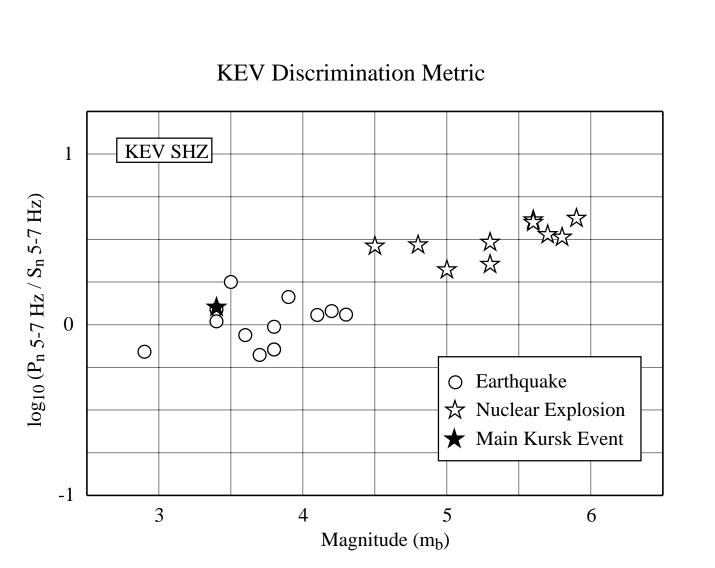


Figure 3a. Broadband KEV record of the main Kursk event. The data have been highpassed at 1 Hz. Amplitudes are normalized to a common scale.

Figure 3b. The ratio of compressional to shear energy for a series of explosions earthquakes recorded at KEV. The main Kursk event falls in with the earthquake population owing to its large shear wave amplitude. Details on the calculation can be found in Hartse [1998].



## Source Complexity

As seen in Figure 2b, many of the Kursk waveforms have strong shear arrivals. The KEV data in particular, Figure 3, have anomalously large shear energy when compared with some previously recorded explosion waveforms.

A second puzzling feature of the Kursk data is that several, but not all, records have dilatational (downward) first motions. Generally waveforms from explosive sources have consistently compressive (upward) first motions.

Presently, it is not clear why the Kursk waveforms show such complexity, however there are several factors that may have contributed:

(1) Asymmetries induced by the positioning of the Kursk as it lay on the seafloor just before exploding.

(2) Energy generated by the impact of the Kursk with seafloor, assuming a simultaneous impact/explosion scenario.

(3) Enhanced mode conversion owing to a surface

(4) Enhanced scattering along source-receiver paths

(5) Substantially different source phenomenology for underwater explosions compared to atmospheric explosions.

**Bubble Pulse Observation** 

Underwater explosions generate bubbles of hot gases which

expand and contract as they rise to the surface. These

oscillations are effective seismic sources and produce

periodicities in the waveforms known as bubble pulses. The

spectral peaks shown in Figure 4a,b are indicative of the

The dominant bubble pulse frequency, fb. of an underwater

explosion is most dependent on explosive yield and depth of

detonation. We observed an f<sub>b</sub> of 1.45 Hz for the main Kursk

event, leading to a yield estimate of 3000-4500 kg TNT if the

detonation occurred at 80-100 m. Note that this frequency, as

well as its overtones, are much too low to be associated with

bubble pulse from the main Kursk event.

water column reverberations.

## Comments on Yield

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Small seismic events

starting in late Sept.

The absolute yield of the main Kursk event is best determined by comparison with previous records of controlled underwater explosions. Such records exist for a series of calibration shots carried out in the Dead Sea in November, 1999. For the largest of the Dead Sea shots, 5,000 kg high explosive at a depth of 70 m, the IMS reported an  $M_1$  of 4.2; this is not significantly different that the  $M_1$  of 4.0 that the IMS reported for the Kursk event. Furthermore, the array station GERES recorded both events, at almost identical distances, and reported m<sub>b</sub> values of 3.3 and 3.4. This yield estimate of ~ 5,000 kg is roughly consistent with the bubble pulse estimate.

An absolute yield estimate for the precursory Kursk event is difficult to obtain, however its yield relative to the main event can be estimated by considering the difference in magnitudes reported by ARCES, 2.2 M<sub>1</sub> and 4.0 M<sub>1</sub>. Using a relation based on a wide range of explosion seismograms [Khalturin, 1998], this magnitude difference of 1.8 units corresponds to roughly 250 times less energy released by the precursory explosion.



from visiting the wreakage.

Post August Seismicity

### Conclusions

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Between August 13 and mid September 2000 there

was no seismicity in region around the Kursk.

However, since September 22 there have been more

than 44 seismic events detected in the region (the

locatons shown to the left were located with PMEL

and the August 12 event as ground truth). The events

are clustered in time: every few days there are 2 to 9

events, and then several days of no events. The

majority of the events are magnitude ML 1.4-1.6,

using depth charges to discourage foreign submarines

In November, CNN reported that the Russians were

although some events are as large as 1.82.

The main Kursk event on 12 August 2000 was seismically recorded at distances of over 5,000 km, and released energy equivalent to 3-7 x 10<sup>3</sup> kg TNT.

- Owing to the clear observation of a bubble pulse, the main event was the direct result of an explosive source and not an impact or collision.

- A precursory event with an essentially identical location to the main event occurred 135 s before the main event. The precursory event had an energy release that was approximately 250 times less than that of the main event.

- The high degree of similarity between the precursory waveforms and the main event waveforms supports the idea that not only were the two events located at nearly the same position but that they had similar source mechanisms as well.

-- It is most likely that the precursory event was a disabling explosion which directly or indirectly led to the catastrophic explosion 135 s later.

Cole, R.H., Underwater Explosions, Princeton Univ. Press, Princeton, New Jersey, 1948.

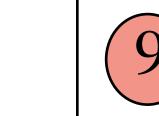
Hartse, H.E., The 16 August 1997 Novaya Zemlya seismic event as viewed from GSN stations KEV and KBS, Seism. Res. Lett., 69, 206-215, 1998.

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Khalturin, V.I., T.G. Rautian, and P.G. Richards, The seismic signal strength of chemical explosions, Bull. Seism. Soc. Am., 88, 1511-1524, 1998.

Wiechecki Vergara, S., H.L. Gray, and W.A. Woodword, Statistical developments in support of CTBT monitoring, DTRA Technical Report, Contract No. DSWA01-98-C-0131, 2000.

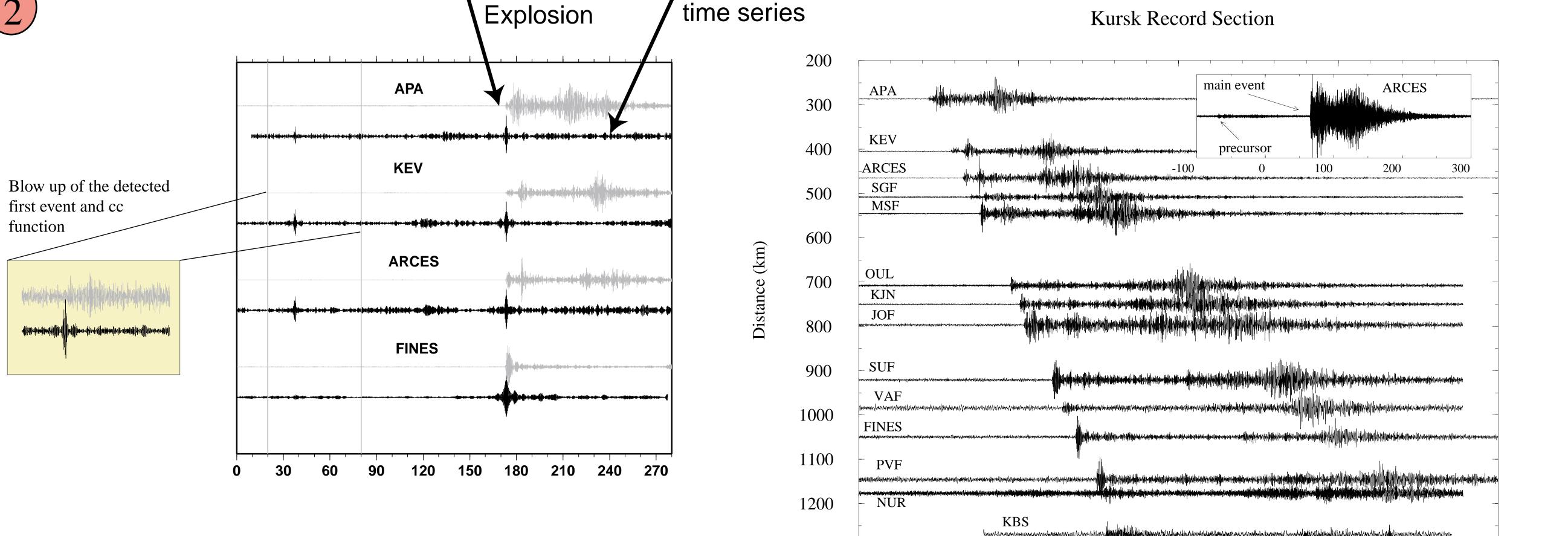
Willis, H.F., Underwater explosions, time interval between successive explosions, British Report WA-47-21, 1941.



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Several colleagues helped in the retrieval

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1300

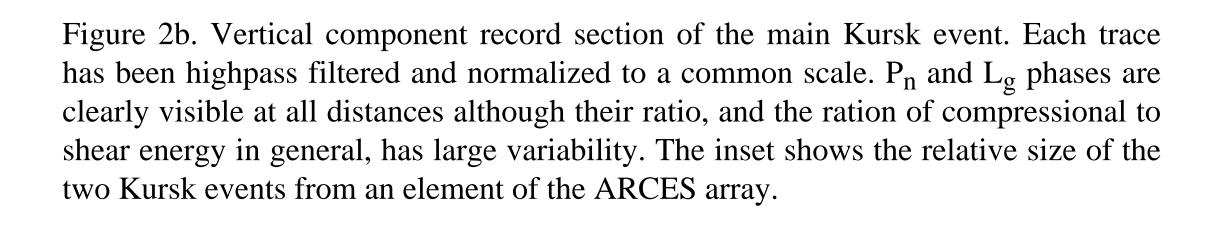
Cross correlation

shown in gray, while the crosscorrelograms are shown in black. A precursory event is clearly visible with at a lag of 135.75 s. The similarity of the precursory waveform with the main waveform is such that the two events must have occurred at nearly identical locations and with similar source mechanisms. We applied the WCD for a time span of +/- 24 hours with respect to the main Kursk event and found no other statistically significant peaks.

Figure 2a. Application of the Waveform Correlation Detector of

Wiechecki Vergara [2000] to the Kursk data. The actual waveforms are

first event and cc



Travel Time (s)

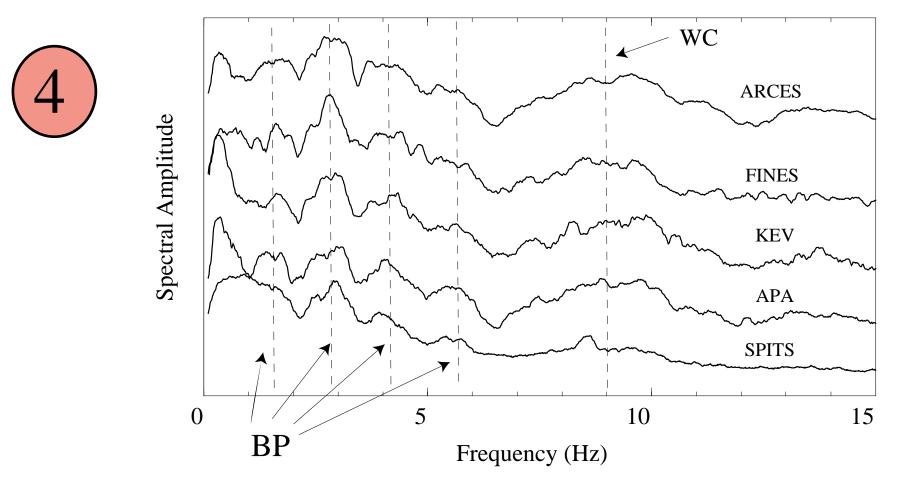


Figure 4a. Spectral amplitudes from several array stations and KEV. Time windows were chosen to bracket the P energy and the spectra have been stacked and smoothed. The spectral peaks indicated by the dashed lines are related to the bubble pulse (BP) oscillations and water column (WC) reverberations.

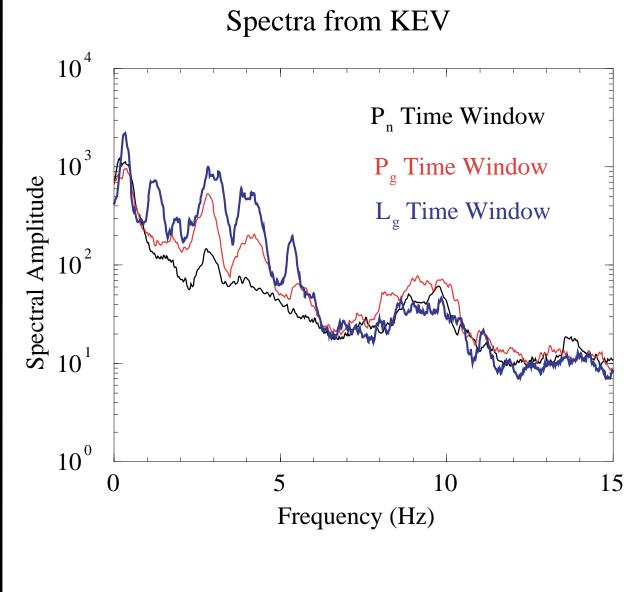


Figure 4b. Smoothed spectra from the vertical component, short period (100 sps) instrument at KEV. Time windows were chosen to bracket three different arrivals: Pn, Pg, and Lg. Peaks and troughs due to the bubble pulse and reverberations in the water column are visible.

# Bubble Pulse Frequency --- Kursk, f<sub>b</sub>=1.45 Hz Detonation Depth (m)

Figure 4c. The analytical dependence of bubble pulse frequency, f<sub>b</sub>, with explosive yield and depth of detonation [Willis, 1941; Cole, 1948]. Lines of constant frequency are shown, and the Kursk fb of 1.45 Hz is indicated by the dashed line.